

NOTES ON BASE

This map sheet is one of a series covering that part of the surface of Mercury that was illuminated during the Mariner 10 mission (for details see Introduction, 1975). The source of map data was the Mariner 10 television experiments (Davies, 1975).

ADOPTED FIGURE

The map projections are based on a sphere with a radius of 2439.4 km.

PROJECTION

The Mercator projection is used for this sheet, with a scale of 1:5,000,000 at the equator. Latitudes are based on the assumption that the spin axis of Mercury is perpendicular to the plane of the orbit. Longitudes are positive westwards in accordance with the usage of the International Astronomical Union (IAU, 1971). Meridians are numbered so that the reference crater named Iku Kal (lat 0.0°, lon 20° W) is centered on long 20° W (Murray and others, 1974; Davies and Batson, 1975).

CONTROL

Planimetric control is provided by photogrammetric triangulation using Mariner 10 pictures (Davies and Batson, 1975). Discrepancies between images in the base mosaic and computed control point positions appear to be less than 2 km.

MAPPING TECHNIQUES

Mapping techniques are similar to those described by Batson (1973a, 1973b). A mosaic was made with pictures that had been digitally transformed to the Mercator projection. Shaded relief was copied from the mosaic and portrayed with uniform illumination with the aid of the Mariner 10 pictures besides those in the base mosaic were examined to improve the portrayal. The shading is not generalized, and may be interpreted with nearly photographic reliability since 1972, Inge and Bridges, 1976).

NOMENCLATURE

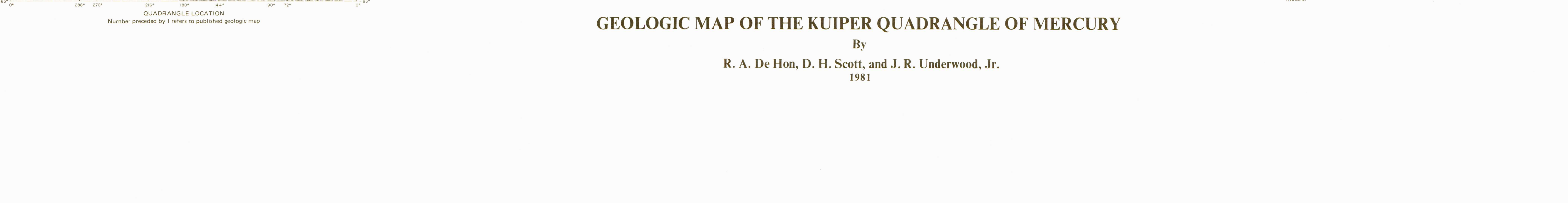
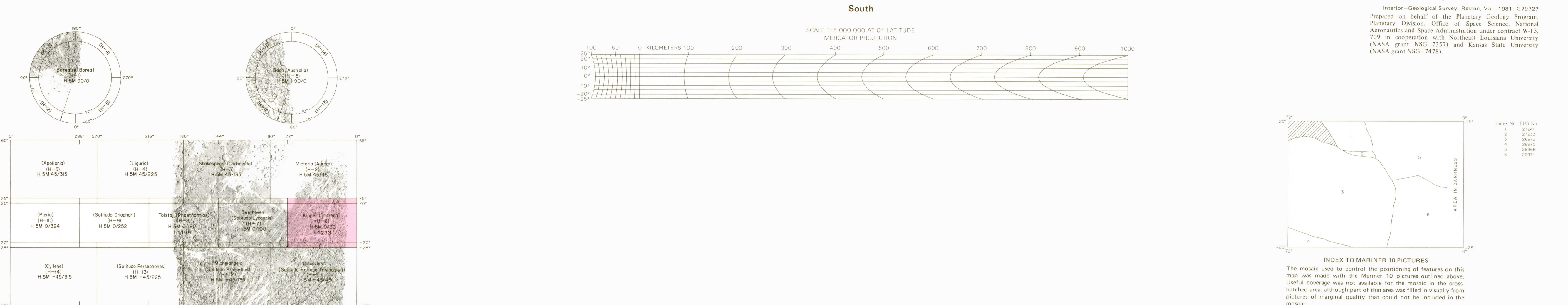
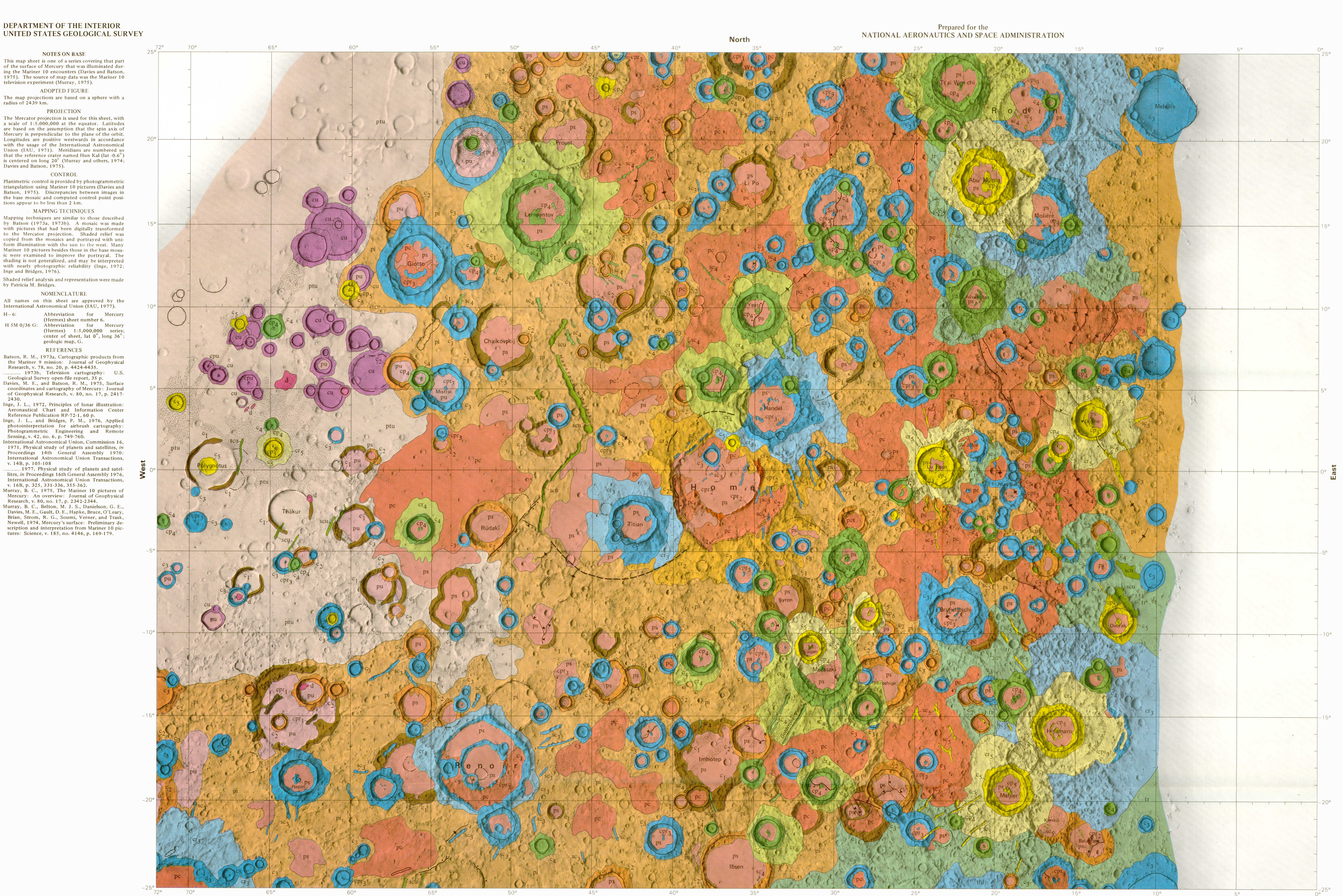
All names on this sheet are approved by the International Astronomical Union (IAU, 1977).

H-6

Abbreviation for Mercury (Hermes) sheet number.
H SM 0/36 G
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Center of sheet, lat 0° lon 36° W.
Geologic map G.

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CORRELATION OF MAP UNITS

PLAINS MATERIALS	TERRA MATERIALS	CRATER AND BASIN MATERIALS
ps	thl	cs
pu	tr	cs2
pc	cu	cs3
pi	cu2	cs4

The Kuiper quadrangle, located in a heavily cratered region of Mercury, includes the young, 55-km-diameter crater Kuiper (11° S, 31° W), which has the highest albedo recorded on the planet (Hapke and others, 1975), and the small crater Iku Kal (0.6° S, 20° W), which is the principal reference crater for nomenclature (Davies and Batson, 1975). Impact craters and basins, their numerous secondary craters, and heavily to lightly cratered plains are the characteristic landforms of the region. At least 10 multiringed basins ranging from 150 km to 440 km in diameter are present. Inasmuch as multiringed basins occur widely and that part of Mercury photographed by Mariner 10, as well as on the Moon and Mars, they offer a potentially valuable basis for comparison between these planetary bodies.

Basic information about the planetary surface of the Kuiper quadrangle is provided by three sequences of high-quality photographs designated Mercury I, II, and III, obtained during the incoming phases of three encounters of the Mariner 10 spacecraft with Mercury. Mercury I includes 75 whole-frame photographs of the Kuiper quadrangle; Mercury II, 13 whole-frame photographs; and Mercury III, 70 quarter-frame photographs. The photographs include 19 exposures in the southern part of the quadrangle. For example, Mariner 10 photography, as Davies and others (1978). The most distant of the photographs was taken at an altitude of 89,879 km, the closest at an altitude of 1,546 km. Resolution, therefore, varies widely but both viewing and solar illumination angles precludes a high degree of mapping consistency. The easternmost 10° of the quadrangle is beyond the evening terminator. A low angle of solar illumination and a high viewing angle make possible discrimination of topographic detail near the terminator. Higher angles of solar illumination and lower viewing angles make it increasingly difficult to discern topographic variations to the west. Many geologic units cannot be specifically identified because of unfavorable viewing geometry west of approximately 55° W. Thus, mapping reliability decreases westward.

Mapping methods and principles are adapted from those developed for lunar photogeologic mapping (Wilhelms, 1970, 1972; Wilhelms and McCauley, 1971). Map units are distinguished on the basis of topography, texture, and albedo and are ranked in relative age on the basis of superposition and function relations. Density of superposed craters, and sharpness of topography. Because of the lack of a widespread, easily identifiable stratigraphic datum on the part of Mercury, a morphologic classification of crater and basin materials was the basis for determining relative ages of many materials. A photostratigraphic map of the best available photogeologic data is provided for interpretation and mapping.

The rock units are subdivided into three major groups: plains materials, terra materials, and crater and basin materials. The plains and smooth terra units are considered to be volcanic in part, and thus may have a different origin from the impact breccias and channeled regolith forming the rough terra and crater deposits.

The oldest rocks exposed in the quadrangle are the intercrater plains material (unit pi) and the rims of the oldest craters and basins. Collectively these rocks form a relatively subdued terrain of moderate relief. They are similar to some of the rolling hills and terraces and hilly and pitted materials in the southern lunar highlands, particularly in the Parthia (Holt, 1974) and Tychus (Paine, 1972) quadrangles. The intercrater plains material is commonly marked by the soft outlines of numerous overlapping secondary craters producing a subdued hummocky texture. It is gradational in places with cratered plains material (unit pc), which forms flat, densely cratered surfaces similar to pre-imbrian plains on the Moon (Wilhelms and McCauley, 1971; Scott, 1972). Although both the cratered and intercrater plains deposits are interpreted to be volcanic, the latter has been highly degraded by repeated impacts over a longer period of time. Mariner 10 surface is probably covered by a relatively thick regolith of reworked impact breccias.

The cratered plains material is related to flat with broad ridges and lobate scarps that in places resemble those of some of the lunar maria. It is difficult to obtain reliable crater counts on this unit because many secondary craters cannot be distinguished from primary craters. Cratered plains materials (unit pc) are divided into two classes: pc1 and pc2. They may represent flows extruded after an initial phase of impact. The albedo of the cratered plains is intermediate compared to that of other materials, but higher than that of the lunar maria, and may reflect lower iron and titanium content (Hapke and others, 1975).

The youngest rock units of the Kuiper quadrangle are the smooth plains materials. Rough terra occurs as overlapping and intermingled ejecta blankets around clusters of large young craters in the eastern part of the quadrangle. The smooth plains material appears to be higher than the rough terra, and the occurrence of dense arrays of fresh secondary craters produces a coarsely textured, hummocky surface at a scale of about 10 km. The effect of roughness is highlighted by the low-solar illumination angle. Ordinarily, rough terra material would be subdivided and mapped as individual ejecta blankets around and belonging to particular craters. In this eastern region, however, the closely grouped craters have about the same age, and it has not been possible to distinguish the boundaries between their aprons in many places.

Smooth plains material covers the floors of numerous craters in all age classifications. Its surface is scored by secondary craters from clusters of c1 and c2 craters in many places in the eastern part of the quadrangle and, within the crater Hon (1° S, 37° W), by secondary craters from the clay craters (Hon (1° S, 42° W) and Hon (4° N, 34° W). Thus the smooth plains unit may have relatively wide age range. Like the cratered plains, it exhibits lobate scarps and low mare-like ridges, but these are generally smaller than those of the cratered plains and more easily resemble the floors of the lunar maria. Although crater counts are more reliable because there are fewer secondaries than in the cratered plains, resolution is a serious constraint to developing crater counts on the relatively coarse smooth plains. Preliminary counts made on a few of the more extensive occurrences of smooth plains show frequencies of about 7.5 x 10⁻⁶ per km² for craters larger than approximately 2.5 km. This frequency is comparable to that of the lunar maria near the Apollo 11 landing site (Greeley and Gault, 1970; Neukum and others, 1975; Meyer and Gault, 1977). Like that of the cratered plains, the albedo of the smooth plains is intermediate compared to other units on Mercury (Hapke and others, 1975) but is relatively high compared to that of the mare basalt on the Moon.

Plains and terra material undivided—Mapped in western part of quadrangle where high solar illumination prevents reliable differentiation between smooth and cratered plains materials (units ps and pu) in craters and other depressions.

INTERCRATER PLAINS MATERIAL—High and highly cratered but intercrater craters are relatively smooth at scale of 5-10 km; most large craters subdivided; many secondaries embayed by gradational with cratered plains (unit pi) in places. Representative area: 8° N, 47° W.

PLAINS AND TERRA MATERIAL UNDIVIDED—Mapped only where resolution of topography is poor because of high-angle solar illumination. Bright rhyolite craters common. Albedo ranges from very high around some craters to intermediate and low. Surface appears generally flat; relief not determinable on photographs. Representative area: 5° S, 62° W. Interpretation: Mostly cratered and intercrater plains materials; may include some smooth plains and rough terra deposits.

HILLY AND LINEATED TERRA MATERIAL—Rugged hills 5-10 km wide and 0.1-1.8 km high, narrow linear troughs, crater rims broken into rings of hills and depressions. Seen only near 25° S, 22° W, maximum development south of Kuiper quadrangle. Interpretation: Structurally modified terrain antipodal to the Caloris Basin in peripheral to large impact basin beyond terminator.

ROUGH TERRA MATERIAL—Very rugged, topographically high and rough surface with numerous large- and medium-size young craters; overlaps and is embayed by smooth plains material. Representative areas: Northeast and southeast corner of quadrangle. Interpretation: Largely ejecta from young large craters that cannot be mapped separately. Highly scored by secondary impacts that formed grooves and chains.

CRATER MATERIAL UNDIVIDED—Materials of craters too poorly resolved for relative age classification.

CRATER MATERIAL—Forms floor, rim, and wall of fresh-appearing craters. Sharp, rugged, complete rim crest; walls typically terraced; well-defined continuous field of relatively crisp secondary craters; bright rays or crater materials superposed on all other craters.

CRATER PEAK MATERIAL—Single or multiple rugged peaks near center of crater.

CRATER RADIAL RIM MATERIAL—Grooved and ridged material extending outward from rim rim as lineated apron with many secondary craters.

SECONDARY CRATER MATERIAL—Shallow, circular to elongate craters, commonly overlapping, aligned in chains or clusters.

CRATER MATERIAL—Similar to c1 craters but less sharp rim crest and slightly modified wall terraces; very few superposed craters.

CRATER PEAK MATERIAL—Same as c1 except occurs within c1 craters.

CRATER RADIAL RIM MATERIAL—Same as c1 except surrounds c1 craters.

SECONDARY CRATER MATERIAL—Same as c1 except secondary from c1 craters.

CRATER MATERIAL—Craters and basins with modified or rounded but intact rim; field of secondary craters present around larger craters but less well-developed than around c1 and c2 craters. Moderate number of superposed craters.

CRATER PEAK OR RING MATERIAL—Central peak or moderately rugged rim of mounds in center of basins.

CRATER RADIAL RIM MATERIAL—Same as around younger craters but discontinuous or less well preserved.

SECONDARY CRATER MATERIAL—Secondary craters surrounding larger c1 craters and basins.

CRATER MATERIAL—Shallow flat-floored craters and basins with low rims. No visible secondaries, many superposed craters.

CRATER PEAK OR RING MATERIAL—Moderately rugged rim of mounds in center of large craters or basins.

CRATER RADIAL RIM MATERIAL—Discontinuous or narrow apron around large c2 basins.

CRATER MATERIALS—Craters or basins with extremely low or discontinuous rims.

CRATER PEAK OR RING MATERIAL—Low, discontinuous, raised rims in center of largest c1 craters.

CRATER PEAK MATERIAL UNDIVIDED—Central peaks and high-albedo material in center of c1 craters.

SECONDARY CRATER MATERIAL UNDIVIDED—Material of allied chains of craters of unknown age.

Symbol in parentheses where concealed by mapped younger deposits.

Contact—Dotted where concealed. West of 50° W, poor illumination results locally in dikes which are approximately located or queried where doubtful.

Fault—Ball and bar on downthrown side.

Ridge in plains material—Symbol on ridge crest.

Gentle rounded scarp—Line marks back-scarp points down slope. May be lava flow front or fault scarp under surficial cover.

Irregular scarp—Line marks top; hachures point down slope.

Lincament, topographic or tectonic.

Depression, rimless.

Crater rim crest—Hachures point down crater wall.

Crater rim crest—Greatly subdued or buried.

Ridge crest of multiringed basins.

Bright ray material—Pumice streaks or halos on and around some c1 craters. Visible only at high-sun angle.

CRATERS AND BASINS

Craters are ubiquitous features of the mercurian surface. For the purpose of mapping, a five-fold morphologic classification of craters (fig. 10 in McCauley and others, 1981) is the basis for determining their relative ages. The youngest craters (c1) have sharp rim crests, textured ejecta blankets, and a well-defined field of secondary craters. Under favorable lighting conditions, the youngest craters exhibit bright rays superposed on all older materials. Older craters have increasingly degraded rims and lower relief and have lost their secondary crater fields. The major differences between mercurian and lunar craters are apparently related to the greater gravitational acceleration and the higher impact velocities on Mercury. Continuous ejecta deposits are less extensive, and secondary craters are more sharply defined and clustered nearer their primary craters. Also, Mercury, accreted secondary craters from prominent crater chains radial to large craters.

Craters within the Kuiper quadrangle increase in complexity as they increase in size from simple bowl-shaped craters to complex craters with central peaks to multiringed basins. Kuiper (11° S, 31° W) is a moderate-size crater with a central peak cluster. Brandebach (9° S, 22° W) exhibits an incomplete ring of peaks; and Rodin (22° N, 18° W) is a well-developed double-ringed basin. These three craters are mercurian counterparts in morphology to the lunar craters Copernicus, Compton (or Antoniadis), and Schrödinger. All craters larger than about 35 km diameter and basins are filled to some extent with plains materials, and exposed rims of partly buried craters within the basins indicate that the fill is about 700 to 1000 m thick (De Hon and Wakom, 1976).

Basins ranging in age from c1 to c5 were formed during the waning stages of high-impact flux when the surface was virtually saturated with craters and basins. Later cratering history records a decreasing impact flux: of craters larger than 50 km diameter, 42 are dated as c1, 19 craters are assigned to c4 and 9 craters are c5. There is also a decrease in the size of the largest crater or basin formed in each age class from c1 to c5.

STRUCTURE

Structural features are sparse or unobserved in this part of Mercury. The Kuiper quadrangle apparently has none of the scars that occur elsewhere on the planet that have been interpreted as high-angle reverse faults (Strom and others, 1975). The most prominent structures are the rings associated with some large craters or basins, faults that transect crater floors, and lobate scarps and ridges in the plains materials. Most of the faults and scarps that transect crater floors clearly delineate crater filling materials standing at different levels, and in at least two craters (19° S, 37° W, 16° N, 30° W), the traces of the faults in the crater walls indicate that the faults have normal displacements. A few faults cut intercrater areas and trend generally northwest to northeast (Scott and others, 1976).

Ridges are broader than many lunar mare ridges and are confined largely to the cratered plains materials. Antoniadis Dorsum, which is a well-developed broad ridge north of the Kuiper quadrangle, is less well developed at its south end and appears to be the quadrangle as an irregular scarp. A number of linear depressions superficially resemble grabens but are chains of overlapping secondary craters, for example, Goldstone Valley (11° S, 32° W) and Hon Valley (5° N, 46° W).

GEOLOGIC HISTORY

The interpretable geologic history of the Kuiper quadrangle is primarily a record of decreasing meteoroid flux during which large craters and basins formed and plains materials were deposited. A decreasing rate of crater production is indicated by progressively fewer craters in each successively younger crater class. Approximately half of the mapped area contains a high density of craters and multiringed basins formed by the intense early bombardment. It is doubtful that any primary crustal material has been preserved without brecciation and reduction by repeated impacts. The present crater population represents only the craters surviving at the end of the stage of highest meteoroid flux. As the impact flux decreased, cratering rates of possible volcanic origin were deposited in broad, low-lying areas, flooding, embaying, or partially burying pre-existing craters. The youngest multiringed basins (Rodin, and the unnamed basin at 15° S, 15° W) formed near the end of this stage (about c3 basins) and did the Caloris Basin on the opposite side of the planet (McCauley and others, 1981; Schaber and McCauley, 1980). Craters that formed still later during the period of low-impact rates were well preserved. During this late stage, smooth plains materials were deposited in basins, craters, and localized low areas and have low crater density. The youngest craters are sharp rimmed with bright rays.

The relatively small size of Mercury, its lack of atmosphere, and the cratered nature of its surface invite comparison with the Moon. The geologic histories of the two bodies are similar. Both surfaces record a decreasing impact flux. The cratering history of Mercury was concurrent with episodes of lava flooding (cratered plains) that may have obliterated some basins and flooded large areas in a manner similar to the mare filling on the Moon. Bright-rayed craters, such as Kuiper, mark the youngest events similar to Copernican craters on the Moon; some dark patches along the west margin of the quadrangle may represent late volcanism.

Some differences between the Moon and that part of Mercury observed in this quadrangle may be more apparent than real. Apparent differences may be the result of resolution of the imaging system, and small viewing and illumination angles that do not allow inspection of the surface under varied conditions. Real differences may be the result of Mercury's size, gravitational field, proximity to the Sun, internal composition and structure, or timing of major volcanic episodes relative to the decrease in impact craters. Surface differences include the preservation of secondary craters around cratered basins, and the absence of recognizable textured and lineated ejecta blankets such as those surrounding the Imbrium and Orientale Basins on the Moon. Possible differences in volcanic features include the absence of widespread dark mare-type deposits, volcanic domes and cones, and smooth ridges. Whereas plains and terra deposits may be distinguished on Mercury, the distinct linear dichotomy of mare and highlands is not present on that half of Mercury observed by Mariner 10.

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